GRA Annual Meeting 2016:

Delineating Recharge Elevation in Headwater Catchments

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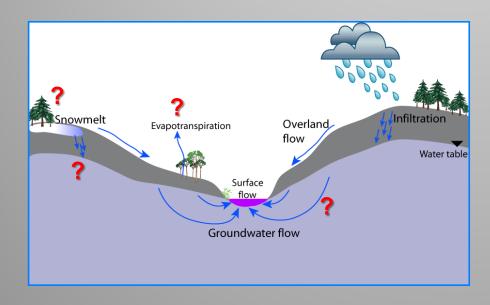






California AB 2480:

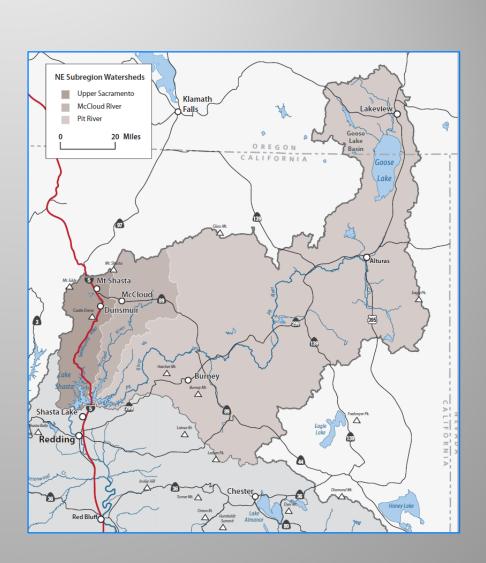
- "...it is state policy to recognize and define source watersheds as integral components of California's water system, and eligible for financing on an equivalent basis with other water infrastructure projects."
- "...(funding for) projects with a demonstrated likelihood of increasing conditions for water and snow attraction, retention, and release under changing climate conditions."





Volcanic, alpine hydrologic setting

- Highly permeable surface materials
- Stream generation from springs
 - Upper Sacramento, McCloud, and Pit River headwaters
 - Springs and wells provide supply for logging operations, towns of Mount Shasta and Weed, bottled water facility, etc.
- ET likely energy (rather than water) limited

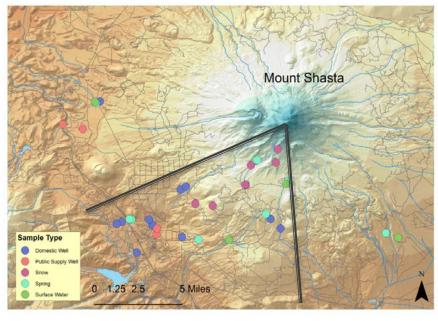


Sampling in headwater areas

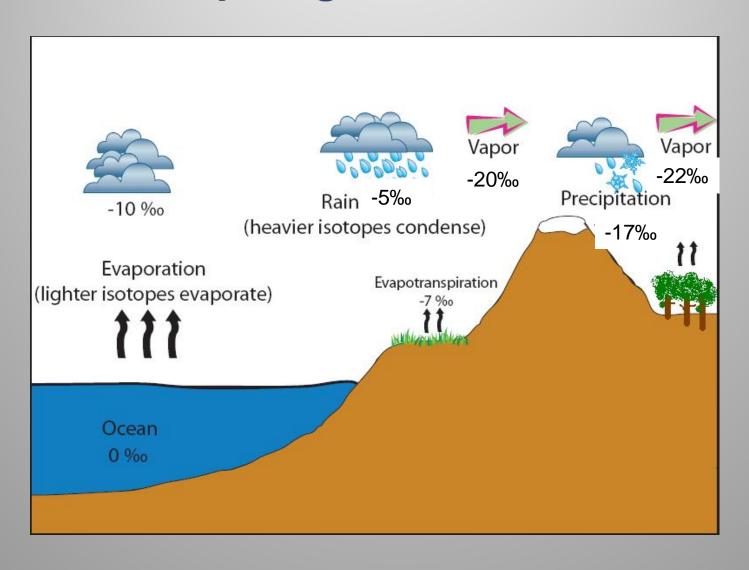
- Limited sampling locations
- Sampled during extreme drought (May and Sept. of 2015)





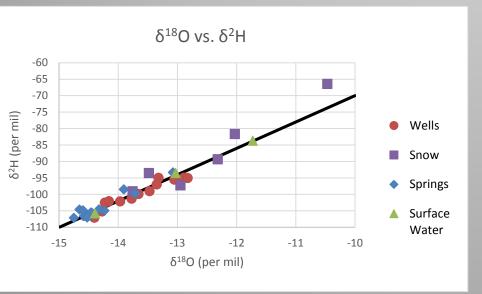


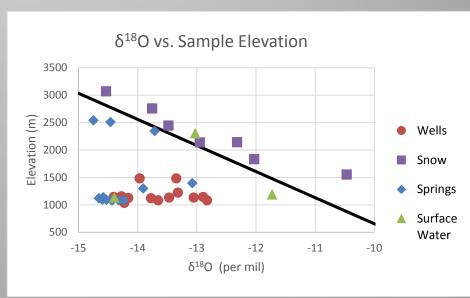
Stable isotope signature: How it works



Stable isotopes: recharge elevation

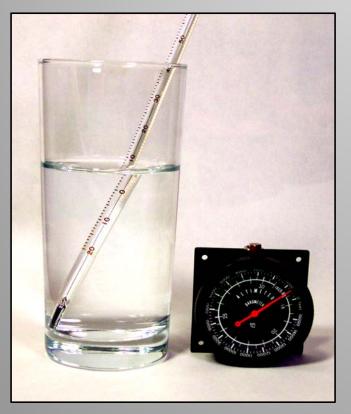
- All samples fall on or near the Global Meteoric Water Line
- No δ¹8O results < -15‰
- Precipitation (snow) follows 'lapse rate' (from Rose et al., 1996)
- Points that fall below the line indicate a source area at higher elevation

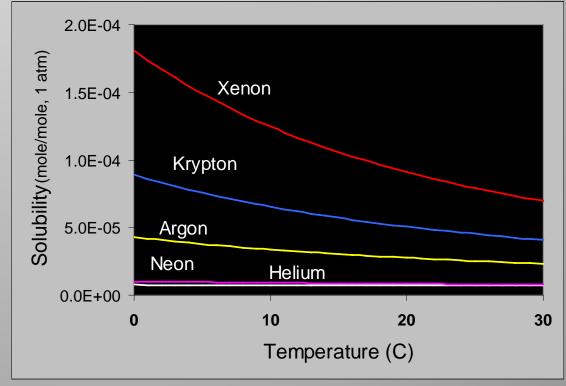




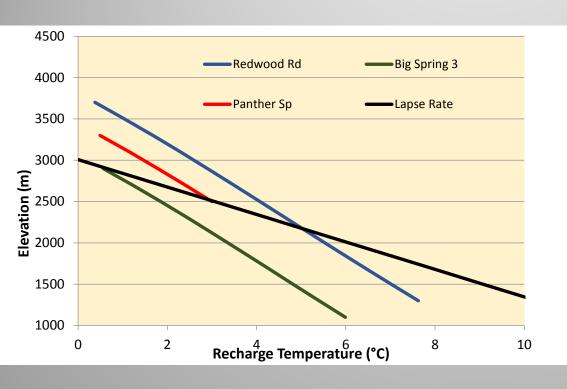
Noble gas signature: How it works

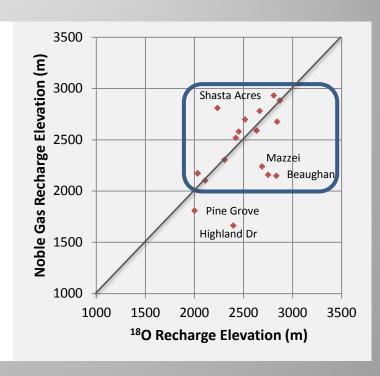
- Solubility depends on temperature and pressure
- Heavier gases have stronger T dependence





Noble gases: recharge elevation

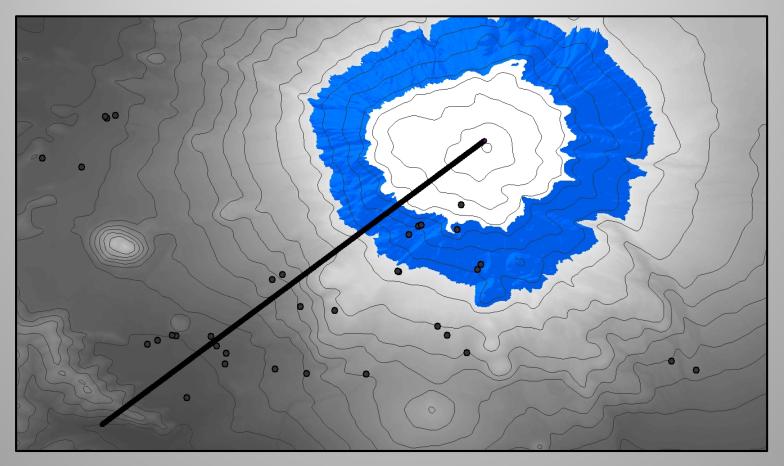




- 3 examples
- Constraints:
 - elevation (top of mountain to sampling elevation)
 - temperature (greater than 0°C, less than discharge temperature)

Good agreement between independent methods

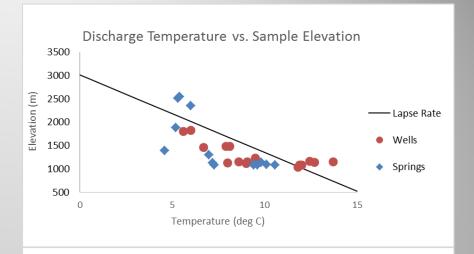
Delineating the recharge area

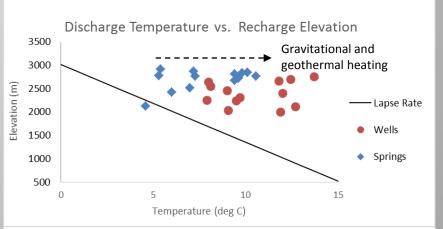


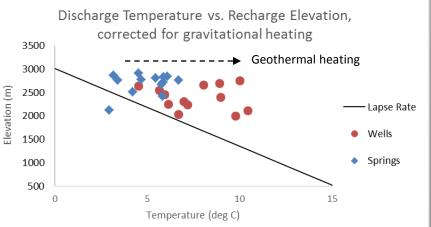
- Elevation range 2200-2800 m
- Do not observe signals from higher or lower elevations

Subsurface Heating

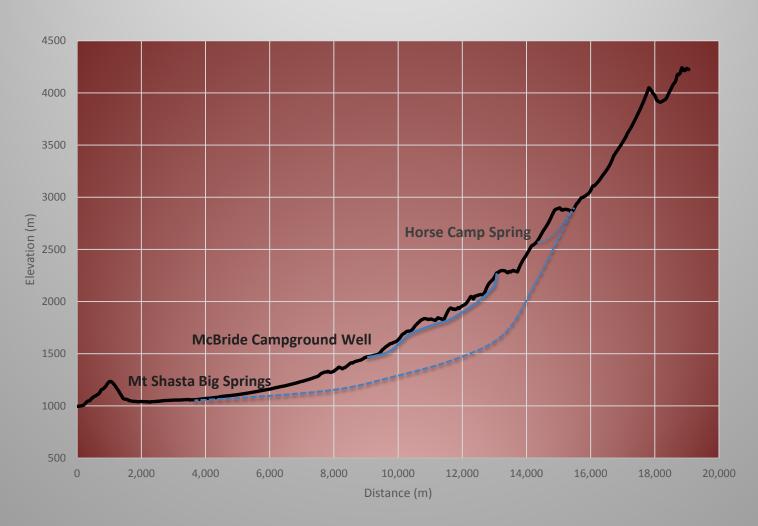
- The rate of change of thermal energy in a parcel of groundwater is the sum of (Manga and Kirchner, 2004):
 - gravitational potential energy dissipation
 - heat transfer to/from the surface by circulating water by conduction (negligible in this setting)
 - geothermal heating
- For a 8°C DT-RT difference and a geothermal gradient of 15°C/km, a maximum flow depth of appx. 500 m is calculated



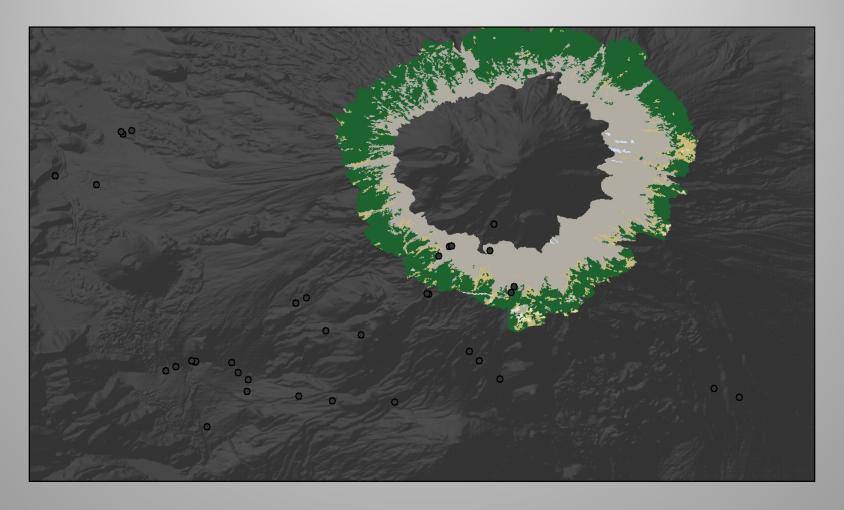




Delineating groundwater flow on Mount Shasta



Land cover in the recharge area



Forest ET likely plays a key role in limiting percolation and recharge

Coverage from: National Land Cover Dataset from the Multi-Resolution Land Characteristic Consortium, www.mrlc.gov, Homer, 2015

There may be a 'sweet spot' in tree coverage

- Too many trees
 - Increased ET
 - Decreased groundwater flow to streams
 - Insufficient openings for snow accumulation

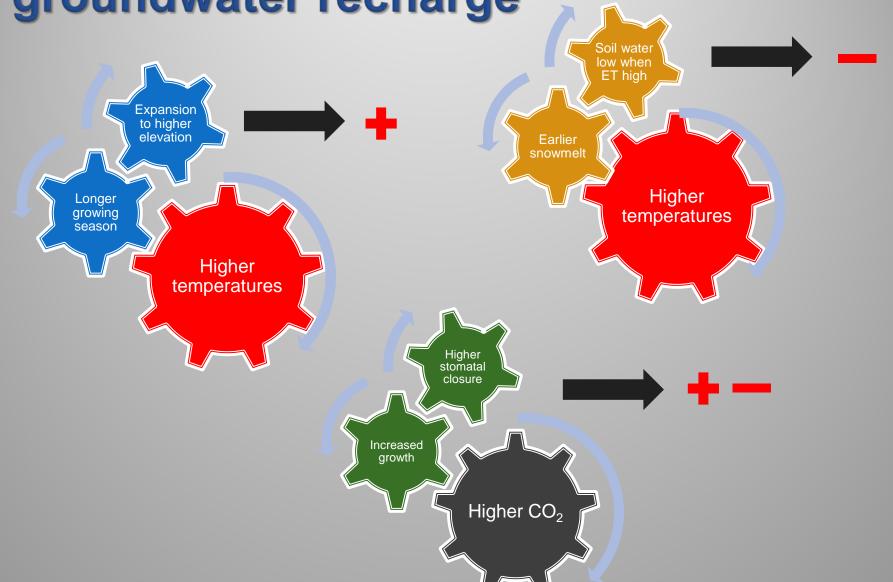


Too few trees

- Increased runoff
- Increased sediment
- Increased stream temperature
- Not enough shade to extend snowmelt season

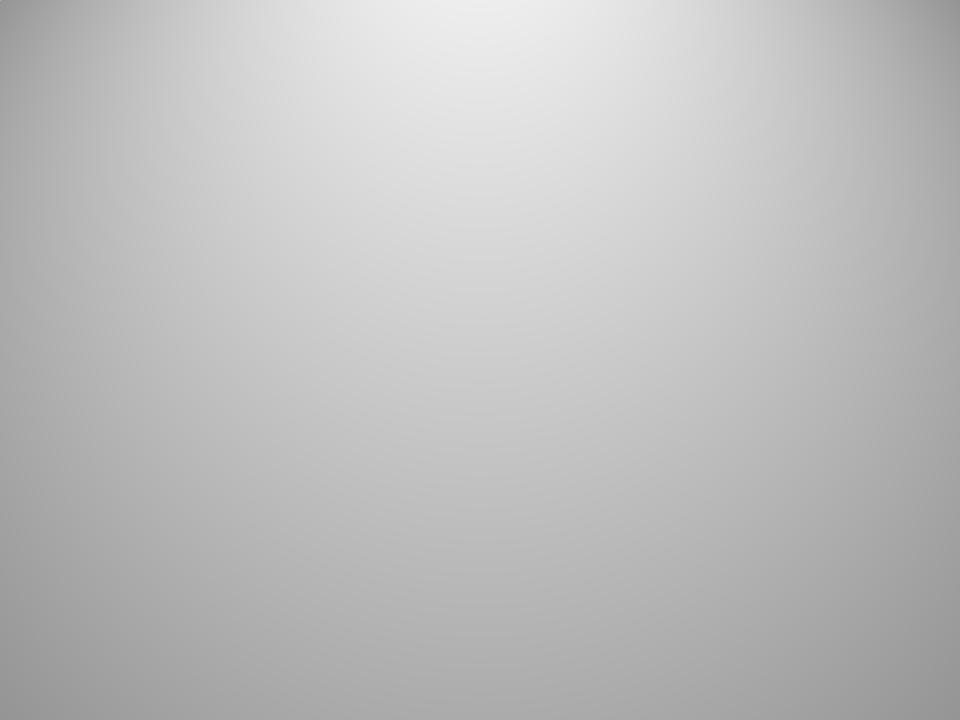


Effects of climate change on ET and groundwater recharge

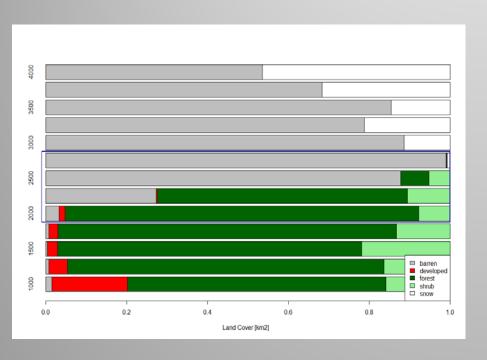


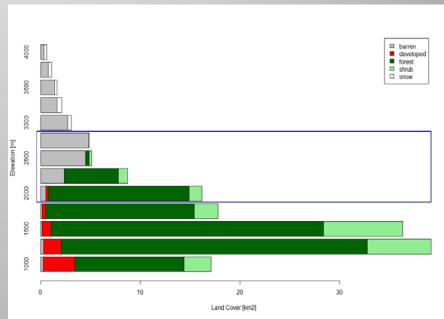
Conclusions

- Headwater areas provide critical late season flow and ecosystem functions
- On Mount Shasta, stable isotopes and noble gas recharge temperatures indicate that recharge occurs predominantly over the elevation range 2200-2900 m (7200 ft to 9500 ft)
- Higher elevations are disproportionately represented because of high precipitation rates and low evapotranspiration over bare ground

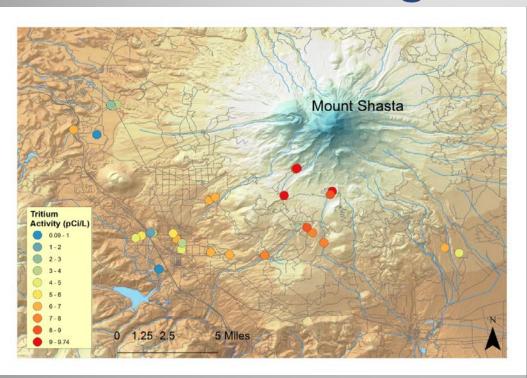


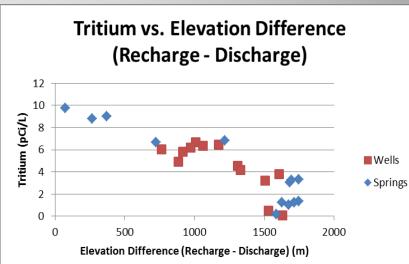
Relationships between elevation and land cover/area

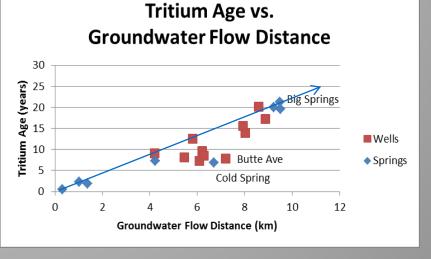




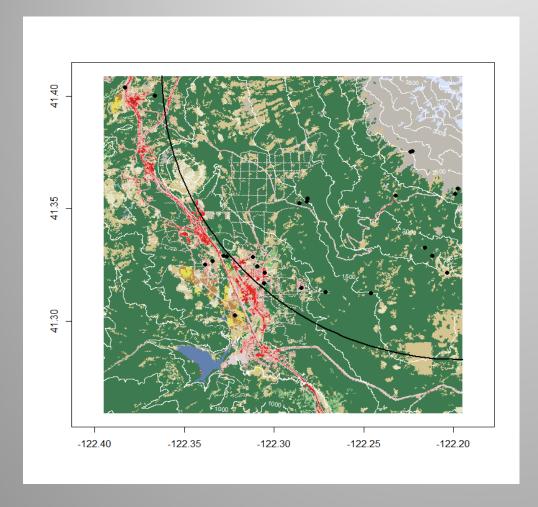
Groundwater ages and flow rates







Land coverage, ET, and recharge



- How much of the melting snow goes to ET instead of infiltrating?
- Changes due to warming climate?